Reflector-preserved lithological upscaling Felix J. Herrmann, Mohammad Maysami*, and Yves Bernabe**

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* Currently at Stanford ** MIT Slogan

"...from seismic reflectivity to connectivity..."

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Motivation

Equivalent media techniques

- wash out singularities
- loose specular reflectivity

Because they average density and constitutive parameters

Wish list

Upscaling techniques that preserve singularities = reflectivity

- link *lithology* to *reflectivity* (e.g. volume fraction shale in sand/shale mixtures)
- provide information on the connectivity
- without oversampling

Two strategies

- Replace linear windowed equivalent medium averaging by equivalent medium averaging based on nonlinear approximations (e.g nonlinear approximations with wavelets based on recent developments in applied Harmonic analysis).
- 2. Use (rock) physical arguments based on the existence of *critical phenomena* in *statistical mechanics* (e.g., phase transitions in percolation theory)

[Schoenberg '88, '89, '92']

Equivalent-medium (EM) approaches

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Wave-equation driven (homogenization)

- anisotropy
- difference (harmonic) averages for density
- static behavior of waves, i.e., the centroid

Mixture-model driven (binary mixtures)

- HS bounds
- Voigt-Reuss

[FJH and Bernabe, '04] [Bernabe and FJH, '04] [Maysami and FJI '08]

Our approach

Include *connectivity* in models for the effective properties of bi-compositional mixtures <=> SWITCH

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- sand-shale, gas-hydrate, opal/opal CT
- upper-mantle mineralogy

Studied two cases:

- elastic properties upper mantle
- fluid-flow properties synthetic rock

[FJH and Bernabe, '04] [Bernabe and FJH, '04]

Approach cont'd

Develop an upscaling methodology based on

- bi-compositional (sand/shale) mixtures
- two litho phases (LP/HP), namely weak and strong
- assume volume fraction (p) increases linearly with depth

Approach cont'd

Model predictions:

 volumetric properties vary smoothly as a function of the volume fractions

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• transport properties may not...





HP Shale

volume fraction





HP Shale

volume fraction

[FJH and Bernabe, '04]

Percolation model



[Knight '05]

Our approach

Incorporate geometry in description binary mixtures.

Distinguish between

 volumetric properties (density & porosity) do not depend on geometry/connectivity

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 transport properties (permeability, stiffness, wavespeed) - depend on geometry/connectivity

For $p < p_c$,

 weak mixture with random disconnected strong inclusions

- with increasing depth more strong inclusions are deposited
- at a critical volume fraction (depth), a
 connected cluster of strong HP is formed

For $p = p_c$,

 an infinite cluster of connected strong (HP) material is formed SLIM 🔶

For $p > p_c$,

not all HP inclusions are part of the infinite HP cluster

- isolated HP clusters are embedded in the remaining LP to form a mixture M
- volume fraction that belongs to HP infinite cluster

$$p^* = p\left(\frac{p-p_c}{1-p_c}\right)^{\beta}$$
 for $\beta > 0$

Switch

Strength of material proportional to cluster size, i.e.,

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$$p^* = \begin{cases} 0 & \text{if } p < p_c \\ p \left(\frac{p - p_c}{1 - p_c}\right)^\beta & \text{if } p \ge p_c. \end{cases}$$

Generates a fractional-order singularity at the critical volume fraction.

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Universality scale exponents

Scale exponents of Percolation depend on

- dimension & type e.g. Bond vs Site Percolation
- statistical properties of the mixing
- if isotropic, then the scale exponents are universal

We use Site Percolation in 3-D yielding $\beta=0.41$

Volume fraction of mixed material is $q^* = 1 - p^*$.

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To model the *mixture*, we need the volume fractions of its LP/HP parts

$$q_M = (1-p)/((1-p) + (p-p^*))$$
 and $p_M = (1-q_M)$,

yielding

$$p_M = 1 - \frac{q}{1 - p\left(\frac{p - p_c}{1 - p_c}\right)^{\beta}}.$$

Ledbetter et al. (1984); Gai et al. (1984); Deptuck et al. (1985); Turosov et al. (1986); Marion and Nur (1989); Favier et al. (1997); Novikov et al. (2001),Stauffer and Aharony (1994),Herrmann and Bernab´e, 2004a; Bernab´e et al., 2004)

Percolation

Well known that *binary* mixtures are strong when strong material is connected and weak otherwise.

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Onsets of connectivity yield transitions similar to phase transitions predicted by Statistical Mechanics

e.g. the onset of magnetization below Cury temperature

Takes connectivity into account...

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[Hashin and Shtrikman (1962)]

Bounds

Both the HP and LP phases are elastically isotropic and HP inclusions are spherical so the HP/LP mixture is locally (statistically) isotropic.

Model materials with *isolated randomly* distributed inclusions inside *connected* matrix.

Use upper bound when strong component forms the connected matrix.

Use lower bound otherwise.

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[Marion and Nur (1989); Favier et al. (1997); Novikov et al. (2001); Saidi et al. (2003)]

Bounds

Bulk modulus **above** critical depth

$$K = K_{LP} \left(1 + \frac{p(K_{HP} - K_{LP})}{q(K_{HP} - K_{LP})a_{LP} + K_{LP}} \right)$$

with p = q - 1 and

 $a_{LP} = 3K_{LP}/(3K_{LP} + 4G_{LP})$

[Marion and Nur (1989); Favier et al. (1997); Novikov et al. (2001)] [Saidi et al. (2003), Herrmann and Bernabe '04; Bernabe and Herrmann, '04]

Bounds

Bulk modulus **below** critical depth

$$K = K_H \left(1 + \frac{q^* (K_M - K_{HP})}{p^* (K_M - K_{HP}) a_{DHLP} + K_{HP}} \right)$$

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with

$$K_M = K_{LP} \left(1 + \frac{p_M (K_{HP} - K_{LP})}{q_M (K_{HP} - K_{LP}) a_{LP} + K_{LP}} \right)$$

Similar expressions hold for shear modulus G.

ρ_1	=	1700	
ρ_2	=	1400	
K_1	=	3.5e9	
K_2	=	2.8e9	
G_1	=	1.5e9	
G_2	=	0.9e9	
n		0 3116	_

$p_c = 0.3116$ $\beta = 0.41$ Equivalent medium



- Profile is smooth
- HS bounds are narrow

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Reuss-Voigt are wide

$ ho_1$	=	1700
$ ho_2$	=	1400
K_1	=	3.5e9
K_2	=	2.8e9
G_1	=	1.5e9
G_2	=	0.9e9
p_c	=	0.3116
ß		0.41

= 0.3116 = 0.41 Percolation model



- Below Pc, HP is disconnected, use lower bound
- Above pc, HP is connected, switch to upper bound with appropriate volume fractions

- Switching leads to
 singularity at p = pc
- Use Reuss-Voigt



Density varies smoothly

Velocity does not

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Singularity generates specular reflectivity





HP Shale

volume fraction





HP Shale

volume fraction

Singularity analysis



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Segmentation & estimation







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Used to constrain the scale exponent for well to seismic tie.

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Opal-Opal CT transition



[Schoenberg '88, '89, '92']

Upscaling problem

How can we upscale to preserve reflectivity?

Mike's proposal is to do a moving average equivalent medium averaging-i.e.,

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$$c_{v,\sigma}(z) = \sqrt{\frac{1}{(\rho * \phi_{\sigma})(z)(\kappa^{-1} * \phi_{\sigma})(z)}}$$
$$c_{h,\sigma}(z) = \sqrt{\frac{(\rho^{-1} * \phi_{\sigma})(z)}{(\kappa^{-1} * \phi_{\sigma})(z)}}$$

[Foldstad and Schoenberg, '92]

Upscaling problem

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... But in that case we need to "oversample" by a factor of 10 and this may lead to difficulties during inversion...

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Upscaling problem

We can use our percolation model instead...

Synthetic log



Velocities



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Cross plots



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Upscaling dillema

Suppose we are given volume fractions p(z) for shale.

Use the Percolation model to compute fine-grained velocities.

Two options to upscale:

- average fine-grained velocities and densities but this smoothes out the switch
- average the volume fractions because this preserves the switch

Upscaled lithology



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Upscaled EM velocities

Equivalent-medium upscaled velocities,



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Upscaled Percolation velocities

Percolation-model upscaled velocities,



Upscales EM reflectivities

Equivalent-medium upscaled reflectivity,



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Upscaled Percolation reflectivities

Percolation-model upscaled reflectivity,



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Observations

At fine scales, the (zero-order) singularities in the lithology dominate the reflectivity

At coarse scales,

- EM-based reflectivity smoothes out
- Percolation-based reflectivity is persistent and is dominated by the (fractional)-order singularity

Morphology?





Morphology?





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Conclusions

Percolation model preserves the singularities

Switch model provided "access" to the finestructure (connectivity) from macroscopic waves

Rigorous mathematical framework for the "shapes" of these percolation-induced transitions is an open problem...

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"...a few words"

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Thank you for your attention!

more information <u>slim.eos.ubc.ca</u>